



DISCUSSION

The Marine geology and geomorphology was mapped in the Offshore of Carpenteria map area from approximate Mean High Water (MHW) to the 100-metric limit of California's State Waters. MHW is defined as the vertical distance of 1.33 m above the North American Vertical Datum of 1988 (NAVD 88) (Webb and others, 2005). Offshore geologic units were delineated on the basis of integrated analyses of adjacent onshore geology with multibeam bathymetry and backscatter imagery (sheets 1, 2, 3), seafloor-sediment and rock samples (Reid and others, 2006), digital camera and video imagery (sheet 6), and high-resolution seismic-reflection profiles (sheet 8).

The onshore geology was compiled from Dibblee (1986), Tan and others (2003a,b), Tan and Clahan (2004), and Minor and others (2009). Unit ages, which are from these sources, reflect local stratigraphic relations.

The offshore part of the map area largely consists of a relatively shallow (less than about 45 m deep), gently sloping shelf (see Fig. 1). The shelf is underlain by sediments deposited from turbidity currents that coastal waters that drain into the inner Mesozoic. Shelf area is primarily sandy (unit **Orms**) at depths less than about 25 m, and at depths greater than about 25 m, are the more fine-grained sediments (very fine sand, silt, and clay) of unit **Orms**. The boundary between units **Orms** and **Orms** is based on observations and extrapolation from sediment sampling (see, for example, Reid and *Others*, 2006) and camera ground-truth surveying (see sheet 6). It is important to note that the boundary between units **Orms** and **Orms** should be considered transitional and approximate and is expected to shift as a result of seasonal-to-annual to decadal-scale cycles in wave climate, sediment supply, and sediment transport.

seasonal to annual-to-decadadal cycles in wave climate, sediment supply, and sediment transport. The latter is the focus of this paper. The sedimentary record of their modern-day sea-floor relief and high backscatter (sheet 3), as well as camera observations (sheet 6) and sampling (Reid and others, 2006; Barnard and others, 2009), are found locally in water depths less than about 15 m, except offshores of Rincon Point where they extend to depths of about 21 m. The largest *Ornitho* deposits are present at the mouths of Rincon Creek and Rincon River, and in the Rincon River estuary. The *Ornitho* deposits are associated with the Rincon River and Rincon Creek lags that armor the seafloor and are relatively resistant to erosion. The sediments may, in part, be relict, having been deposited in shallower marine (or even alluvial?) environments at lower sea levels in the late Pleistocene and Holocene; this seems especially likely for the arcuate lobe of unit *Ornitho* that extends 1,700 m offshore from Rincon Point. The Rincon River and Rincon Creek lags are composed of coarse sand and gravel, and are associated with alluvial fan deposits (Minner and others, 2009) and, thus, may have formed as distal-alluvial fan-de fan-delta facies of that system.

others, 2009) and, thus, may have formed from distal-alluvial or fan-delta facies of that system. A second, more widespread, type of bedrock is composed of coarse-grained, calcareous, and often siliceous Pleistocene and Holocene Peat Formation (unit *OTp*), primarily on the basis of extrapolation from the onshore mapping of Tan and Ochoaiza (2003a,b), Tan and Clahan (2004), and Miner and others (2009), as well as the cross sections of Redin and others (1998, 2004) that are constrained by industry seismic-reflection data and petroleum well logs. Where uncertainty exists, bedrock is mapped as an undivided unit (*OTb*). These rocks are exposed in structural highs that include the Rincon and the San Juan Mountains. The Rincon crest is also overlain by a thin (<1 m) veneer of sediment, recognized on the basis of high backscatter, flat relief, continuity with moderate-to-high-relief bedrock topography, and in some cases high-resolution seismic-reflection data; these areas, which are mapped as composite units *QmTsTn*, *QmTsOTp*, or *QmTsOTb*, are overlain by a thin sediment layer that may be seasonally and/or annually patterned by sediment movement, or longer term climate cycles.

Two offshore anthropogenic units also are present in the map area, each related to offshore hydrocarbon production. The first (unit **af**) consists of coarse artificial fill associated with construction of the Rincon Island petroleum-production facility near the east edge of the map area. The second (unit **pf**) consists of coarse artificial fill mixed with sediment and shell debris, mapped in outcrops surrounding Rincon Island and at the locations of former oil platforms "Heidi," "Hope," "Hazel," and "Hilda" from the Summerland and Carpinteria oil fields (Barnum, 1998). The Monterey Formation is the primary petroleum-source rock in the Santa Barbara Channel, and the Pico Formation is one of the primary petroleum

The Offshore of Carpinteria major area is in the Ventura Basin, in the southern part of the Western Transverse Ranges geologic province, which is north of the California Continental Borderland (Fisher and others, 2009). This province has undergone significant north-south compression since the Miocene, and recent GPS data suggest north-south shortening of about 6 to 10 mm/yr (Larson and Webb, 1992; Donnellan and others, 1993). The active, east-west, striking, north-dipping San Juan Bautista fault (SJB fault) is the primary fault in the area (Fig. 1). The SJB fault is a left-lateral strike-slip fault and is one of the structures on the California Coast Range fault system (e.g., Jackson and Yeats, 1982; Sorlien and others, 2000). This fault system, in aggregate, extends for about 100 km through the Ventura and Santa Barbara Basins and represents an important potential earthquake hazard (see, for example, Fisher and others, 2009).

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^a Between X and Y directions on the same plot, and paper may change size due to atmospheric conditions; therefore, scale and proportions may not be true on plots of this map.

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